



Abstracts

8th European Ostracodologists'
Meeting

Tartu, Estonia, 22-30 July 2015



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Department of Geology, Institute of Ecology and Earth Sciences, University of
Tartu

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8th European Ostracodologists' Meeting

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Contents

Preface	9
AKDEMİR, D., TANYERİ, M., KÜLKÖYLÜOĞLU, O., ALPER, A., DERE, S., YAVUZATMACA, M., YILMAZ, O. & ÖZCAN, O. Ecology, diversity and a/sexual populations of non-marine ostracods in Muğla, Turkey	11
AKİTA, L. G., FRENZEL, P., HABERZETTL, T., KASPER, T., WANG, J. & REİCHERTER, K. Ostracoda as indicators of subaqueous sediment transport – a case study of turbidite and debrite deposits from Tangra Yumco, Tibetan Plateau	12
ALIVERNINI, M., LAI, Z., FRENZEL, P., HABERZETTL, T., MISCHKE, S., PENG, P., WANG, J. & ZHU, L. A late Quaternary lake level curve for Taro Co, Tibetan Plateau, based on ostracod analysis and OSL dating	13
BENNETT, C., BRAND, P., DAVIES, S., KEARSEY, T., MILLWARD, D., SMITHSON, T. & WILLIAMS, M. Repeat colonisation of temporary water-bodies by Early Carboniferous ostracods and bivalves	14
BIESZKE, B., NAMIOTKO, L. & NAMIOTKO, T. Effect of strong electric field (13.5 kV/m, 50 Hz) on life history characteristics of a cosmopolitan non-marine ostracod morphospecies <i>Heterocypris incognuens</i>	15
BÖRNER, N., DE BAERE, B., FRANCOIS, R., JOCHUM, K. P., FRENZEL, P. & SCHWALB, A. Calibration of past environmental conditions based on trace element composition of ostracod shells from the Tibetan Plateau, China	16
CHITNARIN, A. & CRASQUIN, S. Early Devonian ostracods from the Kuan Tang Formation, Satun province, Southern Thailand	17
COHUO, S., MACARIO, L., PÉREZ, L., NAUMANN, K. & SCHWALB, A. Effects of altitudinal gradients in Neotropical ostracod species composition and distribution: an example from north-central Guatemala	18
COHUO, S., MACARIO, L., PÉREZ, L. & SCHWALB, A. Geographical parthenogenesis in northern Neotropical freshwater ostracodes? Understanding the causes using two widely distributed species	20
DANIELOPOL, D. L., NAMIOTKO, T., von GRAFENSTEIN, U., FUHRMANN, R., DECROUY, L., GROSS, M. & PICOT, L. The implementation of taxonomic harmonisation for Candoninae (Ostracoda, Cypridoidea). A heuristic solution for <i>Fabaeformiscandona tricatricosa</i> (Diebel & Pietrzeniuk).	22
DOJEN, C. & GROOS-UFFENORDE, H. Devonian Ostracodes from Morocco (south-west Dra Valley) and the question of the Emsian/Eifelian boundary	24
EWALD, J., FRENZEL, P., PINT, A., SEELIGER, M. & BRÜCKNER, H. Morphological and behavioural observations from culture experiments with <i>Cyprideis torosa</i>	25
FRENZEL, P., ANSORGE, J., DANIEL, T., LORENZ, S., SCHULT, M. & VIEHBERG, F. Ostracoda as palaeoenvironmental proxies in the Holocene of Stralsund, southern Baltic Sea	26
GHAOUACI, S., YAVUZATMACA, M., KÜLKÖYLÜOĞLU, O., AMAROUAYACHE, M. & GHOUZALA, G. Checklist of the living non-marine Ostracoda (Crustacea) of Algeria	27
GLIOZZI, E. & MARCHEGIANO, M. Rose Bengal and Ostracods: the case of the Lake Trasimeno (Umbria, central Italy)	28
GROSSI, F., FARANDA, C., COSENTINO, D., GLIOZZI, E. & BOWRING, S. A. Late Miocene Mediterranean-Paratethys connection: new evidence from the ostracod fauna of the Strymon Basin (northern Greece)	29
HAJEK-TADESSE, V., ILIJANIC, N., MIKO, S. & BAKRAČ, K. Holocene ostracod assemblages and evolution of the shallow freshwater Lake Vrana near Biograd (Croatia)	30
HONG, Y., YASHUARA, M. & IWANTANI, H. Shallow marine ecological degradation in Hong-Kong: a palaeoecological approach using ostracods	31

HORNE, D. J., MARTENS, K., SCHÖN, I. & SMITH, A. J. Taxonomic harmonisation of merged regional datasets of non-marine ostracods: a heuristic approach and its implications for palaeoenvironmental reconstruction	32
IEPURE, S., WYSOCKA, A., SARBU, S. M. & NAMIOTKO, T. Homeomorphy in subterranean Candoninae: Geometric morphometrics of the valve shape and molecular phylogenetic approaches applied for a new species from a chemoautotrophically based Movile Cave ecosystem	33
KOVÁCS, E. & PIPÍK, R. Sublittoral ostracod fauna of the Upper Miocene - Szák Formation, Hungary	34
KRZYMIŃSKA, J. & NAMIOTKO, T. Ostracod and molluscan palaeoassemblages from the Holocene deposits of the Polish part of the Vistula Lagoon, the Baltic Sea	35
KÜLKÖYLÜOĞLU, O. & VEECH, J. A. Estimating co-occurrence assemblages and environmental tolerance of non-marine Ostracoda	36
LI, X. & LIU, W. Environmental changes in Lake Qinghai, NE Qinghai-Tibet Plateau, over the past 32 ka, inferred from ostracod species and their stable isotopes	37
MACARIO L., COHUO S., PEREZ L., VENCES, M. & SCHWALB, A. Genetic diversity on <i>Cyprretta campechensis</i> and <i>Diaphanocypris meridana</i> group in northern Neotropics, new species or cryptic diversity?	38
MACARIO, L., COHUO, S., PÉREZ, L., KUTTEROLF, S., CURTIS, J. & SCHWALB, A. First evidences of Neotropical glacial/interglacial (220-121 ka BP) climate change based on freshwater ostracodes and geochemical indicators from Lake Petén Itzá sediments, Guatemala	39
MARCH, A., HORNE, D. J., HOLMES, J. & LEWIS, S. G. Ostracods from Middle Pleistocene lake sediments at Marks Tey, Essex, UK: Qualitative and quantitative approaches to palaeoenvironmental reconstruction	41
MARCHEGIANO, M., GLIOZZI, E., CESCHIN, S., MAZZINI, I., MAZZA, R. & ARIZTEGUI, D. Living ostracod assemblages of Lake Trasimeno (Umbria, central Italy)	42
MATZKE-KARASZ, R. & SMITH, R. J. Aspects of reproduction with giant sperm in non-marine ostracods	43
MAZZINI, I., RUSCITO, V., GIUSTINI, F., BRILLI, M., SPADONI, M., DI BELLA, L., VOLTAGGIO, M., SADORI, L., PEPE, C., MASI, A. & GIARDINI, M. The coastal evolution of the Tiber delta area during the last 2ky: a micropalaeontological and geochemical study of the Roman imperial Trajan Harbour (Tiber delta, Italy)	44
MAZZINI, I., GLIOZZI, E., COSENTINO, D., KOVACKOVA, M., ATALAR, M., CASTORINA, F. & LO MASTRO, S. Ostracoda from a late Messinian sabkha environment in the central Anatolia Plateau (Çankiri Basin, Turkey)	46
MEIDLA, T. The ostracod assemblage in the mid-Wenlock (Silurian) 'ostracod limestone', Saaremaa Island, Estonia	47
MESCHNER, S. & FRENZEL, P. A new salinity transfer function for the brackish waters of the Wilderness Area, South Africa, based on Ostracoda and Foraminifera	48
MESCHNER, S., FRENZEL, P. & WÜNDSCHE, M. Late Holocene water balance changes in Groenvlei, a Southern Cape coastal lake in South Africa, as indicated by microfossil analysis	49
MESQUITA-JOANES, F., SAVATENALINTON, S. & SUTTAJIT, M. Niche and spatial effects on a highly diverse tropical ostracod metacommunity	50
METTE, W., THIBAUT, N., & KORTE, C. Ecology of benthic microfossils and depositional environments of Late Triassic (Rhaetian) deep neritic deposits in the Northern Calcareous Alps (Austria) – preliminary results.	51
MEYER, J., WROZYNA, C. & PILLER, W. E. Biogeographical differences in stable oxygen and carbon isotopes of <i>Cytheridella</i> in the Neotropics: the case of the Florida area	52

MICHELSON, A. V., BRADY, K., ASH, J. L., WAMSLEY, K., SPERGEL, J. & PARK BOUSH, L. Extending the reach of precise paleoenvironmental reconstructions into deep time using community-wide trait distributions of ostracods	53
NAMIOTKO, T., MEISSNER, W. & NAMIOTKO, L. Ostracoda of shallow floodplain water bodies in the lower reaches of the Ob River in the taiga/forest-tundra transition zone of the Western Siberian Lowland, Russia	54
NAZİK, A., ÇAPKINOĞLU, Ş., OLEMPSKA, E., ÖZGÜL, N. & ŞEKER, E. Ludlow (Silurian) and Givetian (Devonian) ostracods and conodonts from the İstanbul Zone (Kartal and Tuzla Peninsula), NW Anatolia	55
OLEMPSKA, E. & WACEY, D. Ambient Inclusion Trails in Palaeozoic arthropods (Phosphatocopina and Ostracoda)	56
OLSZEWSKI, P., SELL, J. & NAMIOTKO, T. Ostracods meet bacteria: Species-specific microbiome of freshwater ostracods	57
ÖZCAN, G., KÜLKÖYLÜOĞLU, O., YAVUZATMACA, M., YILMAZ, O., TANYERI, M., AKDEMİR, D., ÇELEN, E., DERE, Ş., DALKIRAN, N. & ALPER, A. Ecology and species diversity of Ostracoda (Crustacea) in Ağrı region (East of Turkey)	58
PARK BOUSH, L., V. MICHELSON, A. & MYRBO, A. Ostracode Distribution in Lakes in the Bahamas as a Response to Sea Level and Climate Change	59
PERRIER, V., WILLIAMS, M., SIVETER, D. J., GOODALL, R., MIKHAILOVA, E., TARASENKO, A., SALIMOVA, F. & KIM, I. A. Quantifying the origins of a pelagic lifestyle in ostracods	60
PIERI, V., ALFONSO, G., MARRONE, F., STOCH, F. & ROSSETTI, G. Distribution of Recent ostracods in inland waters of the Mediterranean area (Greece, Southern Italy, and Malta)	61
PINT, A. & FRENZEL, P. Ostracod fauna associated with <i>Cyprideis torosa</i> – an overview	62
PINT, A., SCHNEIDER, H., FRENZEL, P., HORNE, D. J. & VIEHBERG, F. Late Quaternary lake history of the Siebleber Senke (Thuringia, Central Germany) – methods of palaeoenvironmental analysis using Ostracoda	63
QIN, Y., ZHANG, G. & GU, Y. Ostracod ecology and response to human activities in lakes of the middle and lower Yangtze River plain	64
RODRIGUEZ-LAZARO, J., MARTÍN, M., ANADÓN, P., BARRÓN, E., ROBLES, F., UTRILLA, R. & VÁZQUEZ, A. A Miocene saline lake evolution: Ostracods from Moneva (Ebro Basin, Spain)	65
RYCHLIŃSKA, J., SELL, J. & NAMIOTKO, T. (Un)expectedly high genetic diversity of <i>Heterocypris incongruens</i> (Ostracoda, Cyprididae) from Iberian all-female populations	66
ŞAFAK, Ü. Environmental properties and micropalaeontological investigation of tertiary sequences in Çorlu-Muratlı-Lüleburgaz-Babaeski (Southeastern Thrace, Turkey)	67
SEKO, M. & PIPIK, R. Langhian (middle Miocene) ostracod assemblage from the Carpathian Foredeep	68
SIVETER, D. J., PERRIER, V. & WILLIAMS, M. British Upper Silurian Myodocopes: a new stratigraphical tool for regional and interregional correlation	69
SIVETER, David J., BRIGGS, D. E. G., SIVETER, Derek J. & SUTTON, M. D. A Silurian pentastomid parasitic on ostracods	70
SMITH, D., WILKINSON, I., WILLIAMS, M., ZALASIEWICZ, J. & SCARBOROUGH, J. The landscape of a bronze age riparian community at Wittlesey Cambridgeshire, UK microfaunal applications	71
SPADI, M. & GLIOZZI, E. Redefinition of the Genus <i>Caspiocypris</i> Mandelstam, 1956 (Ostracoda, Candoninae) and its distribution in the Neogene and Quaternary of Italy	72

SÝKOROVÁ, M., PIPÍK, R., LÁNCZOS, T., STAREK, D. & ŠURKA, J. Ecology of living Ostracoda from travertine springs and lakes of Western Carpathians	73
TANAKA, H. Mating behaviour and male upper lip morphology of the genus <i>Parapolycope</i> (Cladocopina): its significance for speciation	74
TANYERİ, M., YILMAZ, O. & KÜLKÖYLÜOĞLU, O. Seasonal distribution and species succession of Ostracoda in Taşlyayla-Seben reservoir (Bolu,Turkey)	75
TRUUVER, K. & MEIDLA, T. Response of ostracods of the Baltoscandian Palaeobasin to the Hirnantian glaciation	76
TUNCER, A., TUNOĞLU, C., DALGÖĞÜSOĞLU, M. K. & AŞKIM GÜMÜŞ, B. Distribution of ostracod assemblages in Çiğdem and Terzili Ponds, Kastamonu, Northern Turkey	77
TUNCER, A., TUNOĞLU, C., KAYSERİ-ÖZER, M. S., AKGÜN, F., ŞEN, Ş. & KARADENİZLİ, L. Paleoenvironmental interpretations and age constraints on Akkaşdağı Formation using ostracods and palynofloras, Çankırı-Çorum Basin, Central Anatolia	78
TUNOĞLU, C., TUNCER, A., AKBULUT, A., GÜMÜŞ, H., KÖSE, T. & ŞALIŞ K. Distribution of ostracod and diatom assemblages in Beyler Dam Pond, Kastamonu, Northern Turkey	79
TUNOĞLU, C., TUNCER, A., SOLAK, C. N., FETHİ, F. Y., PALAS, S. & İLERİ, Ö. Preliminary results on ostracod and diatom assemblages of Lake Eğirdir, Isparta, Western Turkey	80
UFFENORDE, H. Living and Quaternary Ostracoda from the Eastern Adriatic Sea: Biocoenoses, thanatocoenoses or palaeoethanatoenoses?	81
WILKINSON, I. P. The Anthropocene: Ostracods meet Man	83
YASUHARA, M. & DANOVARO, R. Temperature impacts on deep-sea biodiversity	84
YASUHARA, M., HUNT, G., OKAHASHI, H. & BRANDÃO, S. N. Taxonomy of deep-sea trachyleberidid, thaerocytherid, and hemicytherid genera (Ostracoda)	85
YAVUZATMACA, M., KÜLKÖYLÜOĞLU, O. & SARI, N. Comparison of the hemipenis of the genus <i>Heterocypris</i>: a case study for <i>Heterocypris incongruens</i> (Ramdohr, 1808)	86
YAVUZATMACA, M., KÜLKÖYLÜOĞLU, O., AKDEMİR, D., TANYERİ, M., YILMAZ, O., DALKIRAN, N. & ÇELEN, E. On the relationship between the occurrence of ostracod species and elevation in Sakarya region, Turkey	87
YILMAZ, O., KÜLKÖYLÜOĞLU, O., TUNOĞLU, C., NAZİK, A., AKDEMİR, D., YAVUZATMACA, M. & TUNCER, A. Geographical and stratigraphical distribution of the genus <i>Zonocypris</i> MÜLLER, 1898 in Turkey and in the World	87
ZAZZALI, S. & CRASQUIN, S. Ostracods at the Middle-Upper Permian boundary	89
ZENINA, M. A., SCHORNIKOV, E. I. & YANINA, T. A. Specific ostracod fauna of the chocolate-colored clays in North Caspian region	90

APPENDIX (added in the pdf edition)

SAVATENALINTON, S. A new genus and five new species of subfamily Cypridopsinae Kaufmann, 1900 (Crustacea: Ostracoda) from Thailand	91
SAVATENALINTON, S. & SUTTAJIT, M. Two new ostracods (Crustacea: Ostracoda) from Thailand	92

Preface

The ostracod workers are meeting regularly in different places of Europe but this is the first time when the meeting takes place in Estonia. Tartu, the location of the 8th European Ostracodologists' Meeting was decided during the 7th EOM in Graz, Austria, in the summer of 2011. The meeting is hosted by the Department of Geology of the University of Tartu.

The meeting is held in July 22-30, 2015 and the period is divided into three parts. The pre-conference field trip starts from Tallinn on July 22th and takes a small group of people to a number of sites related to the ostracod studies in Estonia. The scientific sessions in Tartu are held from July 24th to 27th, with the mid-conference excursion to the Endla Nature Reserve and inter-drumlin Lake Saadjärv. The post-conference excursion visits the Ordovician and Silurian sections on the Island of Saaremaa and in mainland Estonia, it departs from Tartu on July 28th and terminates in Tallinn on July 30th.

The present abstract volume was prepared for the meeting. 40 talks and 34 poster presentations of this meeting summarize recent advances in ostracod studies, covering a wide range of topics from biology to geoarchaeology. Several business meetings are held during conference.

The organizers thank all contributors and members of the scientific committee and acknowledge financial support from the University of Tartu.

Tõnu Meidla and Oive Tinn

On behalf of the Organizing Committee

Paleoenvironmental interpretations and age constraints on Akkaşdağı Formation using ostracods and palynofloras, Çankırı-Çorum Basin, Central Anatolia

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The study area is located at the southern part of the Çankırı-Çorum Basin, filled by a thick Paleogene-Quaternary sedimentary sequence. The Çankırı-Çorum Basin is one of the most important Cenozoic basin of Central Anatolia. The Akkaşdağı formation, located in the southern part of the basin, unconformably overlies the basement rocks and, in turn, is unconformably overlain by Quaternary alluvial deposits. It is mainly composed of massive mudstones, laminated claystone, gravelly sandstones, bedded limestones, gypsum rose and tuffs. Nine samples have been collected for ostracod analysis. Nine ostracod taxa belonging to eight genera have been determined: *Candona* cf. *C. devexa*, *Candona* sp., *Ilyocypris bradyi*, *Cypris pubera*, *Herpetocypris* cf. *H. chevreuxi*, *Heterocypris salina*, *Potamocypris zschokkei*, *Zonocypris membranae* and *Cyprideis sublittoralis*. One hundred and eighteen samples have also been collected for palynofloral analysis. Coal bearing sediments in Kırıkkale and its vicinity (Central Turkey) yielded palynofloras of the Late Miocene and Early Pliocene age. The Late Miocene palynoflora is characterized by low diversity and abundance of spores and pollens. The Early Pliocene palynoflora, on the contrary, is rather different for the presence of rich and various palynomorphs. Spore and pollen distribution of the Early Pliocene consists of abundant herbaceous and shrubs elements. The paleovegetation underwent significant changes from the Late Miocene to Early Pliocene probably due to changes in temperature and precipitation. Chronostratigraphic ranges of the known ostracod species obtained in this study indicates Late Miocene. The mammalian fauna representing MN12 zone and the radiometric ages gathered from tuffs ($7,1 \pm 0,1$ Ma) in previous studies also point to a Late Miocene age (Messinian). Moreover, identified palynofloras indicate Late Miocene (Messinian) and Early Pliocene (Zanclean). By combining all these data, the age of the formation is suggested as Late Miocene–Early Pliocene. *Candona* cf. *C. devexa*, *Ilyocypris bradyi*, *Cypris pubera*, *Herpetocypris* cf. *H. chevreuxi*, and *Potamocypris zschokkei* were mainly recovered from the lower part of the formation and point to shallow, stagnant and/or slow flowing freshwater to oligohaline water environments. On the other hand, some species (*Heterocypris salina*, *Zonocypris membranae*, *Cyprideis sublittoralis*), mainly observed in upper parts, are known as halophilic that can tolerate mesohaline salinity ranges. *Cyprideis sublittoralis* is dominant in the uppermost part of the formation and according to its population structure (juvenile/adult and valve/carapace ratios), it can be suggested that low energy conditions prevailed along the deposition of the uppermost part of the succession.